LES of Turbulent Flows: Lecture 23

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Fall 2016



1 Lagrangian Particle Dispersion Modeling in LES



Lagrangian Particle Dispersion Modeling in LES

• This is a special lecture on Lagrangian particle dispersion in LES created by Brian Bailey





Lagrangian Particle Dispersion Modeling in LES

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November 18, 2014

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Visualizations



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Lagrangian vs Eulerian Reference Frames

Eulerian





Best for smoothly varying scalar fields (i.e., continuum)

Governing Equation

$$\frac{\partial C}{\partial t} + \frac{\partial u_j C}{\partial x_j} = D \frac{\partial^2 C}{\partial x_j x_j}$$



Lagrangian vs Eulerian Reference Frames





Best for smoothly varying scalar fields (i.e., continuum)

Best for discrete sources, or when details of individual particles are of interest

Lagrangian

Governing Equation $\frac{\partial C}{\partial t} + \frac{\partial u_j C}{\partial x_i} = D \frac{\partial^2 C}{\partial x_i x_i}$

Governing Equation

 $\frac{\mathrm{d}x_i}{\mathrm{d}t} = u_i$

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Numerical solution

$$\frac{\mathrm{d}x_i}{\mathrm{d}t} = u_i$$

$$\frac{x_i(t + \Delta t) - x_i(t)}{\Delta t} = u_i(t)$$

$$x_i(t + \Delta t) = x_i(t) + u_i(t)\Delta t$$



Note

Side Note:

This form assumes particles are massless.

Could add generic velocity (say u_i^*) to account for gravitational settling, inertia, etc.

$$\frac{\mathrm{d}x_i}{\mathrm{d}t} = u_i + u_i^*$$



Numerical solution example

$$x_i(t + \Delta t) = x_i(t) + u_i(t)\Delta t$$

Consider
$$x(0) = 0$$

 $u(x = 0) = 1$, $u(x = 0.5) = 2$, $u(x = 1) = 1.5$
 $\Delta t = 0.1$



Numerical solution example

$$x_i(t + \Delta t) = x_i(t) + u_i(t)\Delta t$$

Consider
$$x(0) = 0$$

 $u(x = 0) = 1, u(x = 0.5) = 2, u(x = 1) = 1.5$
 $\Delta t = 0.1$
 $\underbrace{t \quad x \quad u}_{00}$
 $0.1 \quad 0.1 \quad 1.2$
 $0.2 \quad 0.22 \quad 1.44$
 $0.3 \quad 0.36 \quad 1.73$
 $0.4 \quad 0.54 \quad 0.71$



What's the problem if we want to apply this to LES?



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What's the problem if we want to apply this to LES?



We don't know $u_{s,i}!$



Could neglect it
$$(u_{s,i} = 0)$$

e.g., **Pure Convection:**

Gopalakrishnan, S. G., and R. Avissar, 2000: An LES study of the impacts of land surface heterogeneity on dispersion in the convective boundary layer. *J. Atmos. Sci.*, **57**, 352–371. Near-Canopy Flow:

Bailey, B. N., R. Stoll, E. R. Pardyjak, and W. F. Mahaffee, 2014: The effect of canopy architecture and the structure of turbulence on particle dispersion. *Atmos. Env.*, **95**, 480–489.



Modeling $u_{s,i}$: where should we start?



Modeling $u_{s,i}$: where should we start?

Let's copy the RANS people.

Why? RANS is essentially LES with the grid scale equal to the domain size....so this *should* be easier.



Lagrangian dispersion in RANS:





RANS models



Analogy to molecular motion (Brownian motion): Langevin Equation



 u_i - molecule velocity

 $\mathrm{d}\xi_i$ - random Gaussian process with mean zero and variance $\mathrm{d}t$

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Application to isotropic turbulence: $u_i \rightarrow \text{Lagrangian particle velocity}$

$$\mathrm{d}u_i = -\boldsymbol{a}u_i \mathrm{d}t + \boldsymbol{b}\mathrm{d}\xi_i$$

How do we get a and b?



Langevin Equation: finding b

$$\mathrm{d}u_i = -\boldsymbol{a}u_i \mathrm{d}t + \boldsymbol{b}\mathrm{d}\xi_i \tag{1}$$

b comes directly from Kolmogorov's second hypothesis Lagrangian structure function:

$$D(\Delta t) = \langle (\Delta w)^2 \rangle = C_0 \varepsilon \Delta t$$

Provided Δt is in the internal subrange (i.e., $\tau_\eta \ll \Delta t \ll \tau_L$)



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$$D(\Delta t) = \langle (\Delta w)^2 \rangle = C_0 \varepsilon \Delta t$$

Provided Δt is in the internal subrange (i.e., $\tau_{\eta} \ll \Delta t \ll \tau_L$) Square Eq.1 and take ensemble average:

$$\langle (\Delta w)^2 \rangle = -\langle aw^2 (\Delta t)^2 \rangle = ab \langle w\Delta \xi \rangle \Delta t + b^2 \langle (\Delta \xi)^2 \rangle \Delta t$$
$$\langle (\Delta w)^2 \rangle = b^2 \Delta t = C_0 \varepsilon \Delta t \rightarrow b = (C_0 \varepsilon)^{-1/2}$$



Langevin Equation: finding a

$$\mathrm{d}u_i = -\boldsymbol{a}u_i \mathrm{d}t + \boldsymbol{b}\mathrm{d}\xi_i$$

Using stochastic calculus, we can solve this equation analytically

$$w(t) = w(0)e^{-at} + be^{-at} \int_0^t e^{as}\xi(s) \mathrm{d}s$$



Langevin Equation: finding a

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Using stochastic calculus, we can solve this equation analytically

$$w(t) = w(0)e^{-at} + be^{-at} \int_0^t e^{as}\xi(s) \mathrm{d}s$$

Square this equation and take ensemble average:

$$\langle w^2(t)\rangle = \langle w^2(0)\rangle e^{-2at} + \langle \psi(0)\rangle e^{-2at} \int_0^t e^{as}\xi(s)\mathrm{d}s + \langle \mathbf{b}^2 e^{-2at} \left[\int_0^t e^{as}\xi(s)\mathrm{d}s\right]^2 \rangle$$

$$\langle w^2(t)\rangle = \langle w^2(0)\rangle e^{-2at} + \frac{b^2}{2a} \left[1 - e^{-2at}\right]$$



Langevin Equation: finding a

$$\langle w^2(t) \rangle = \langle w^2(0) \rangle e^{-2at} + \frac{b^2}{2a} \left[1 - e^{-2at} \right]$$
(2)

For homogeneous and isotropic turbulence,

$$\langle w^2(t)\rangle = \langle w^2(0)\rangle = \sigma_w^2$$
 (const.)

Make this substitution and evaluate Eq. 2 at $t \to \infty$

$$\sigma_w^2 = \frac{\boldsymbol{b}^2}{2\boldsymbol{a}}$$

$$\boxed{\bm{a} = \frac{\bm{b}^2}{2\sigma_w^2} = \frac{C_0\varepsilon}{2\sigma_w^2}}$$



Application to homogeneous isotropic turbulence

$$\mathrm{d}u_i = -\frac{C_0\varepsilon}{2\sigma^2}u_i\mathrm{d}t + (C_0\varepsilon)^{1/2}\,\mathrm{d}\xi_i$$

for homogeneous isotropic turbulence, $\frac{2\sigma^2}{C_0\varepsilon} = \tau_L \text{ is the integral timescale}$ $\mathrm{d}u_i = -\frac{u_i}{\tau_L}\mathrm{d}t + (C_0\varepsilon)^{1/2}\,\mathrm{d}\xi_i$





I Gives correct integral timescale of τ_L (long-time behavior)



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$$\mathrm{d}w = -\underbrace{\frac{w}{\tau_L}}_{\mathbf{I}} \mathrm{d}t + \underbrace{(C_0\varepsilon)^{1/2}\,\mathrm{d}\xi_i}_{\mathbf{I}}$$

I Gives correct integral timescale of τ_L (long-time behavior)
 II Makes velocity consistent with Kolmogorov's second hypothesis (short-time behavior)

Inhomogeneous Turbulence (in 1D)

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Inhomogeneous Turbulence (in 1D)

Well-Mixed Condition

Well-Mixed Condition¹ or Thermodynamic Constraint²

An initially well-mixed (uniform) particle distribution must remain well-mixed for all time in the absence of sources or sinks (second law of thermodynamics).

¹Thomson, D. J., 1987: Criteria for the selection of stochastic models of particle trajectories in turbulent flows. *J. Fluid Mech.*, **180**, 529–556.

²Pope, S. B., 1987: Consistency conditions for random walk models of turbulent dispersion. *Phys. Fluids*, **30**, 2374–2379.

Langevin Equation: Inhomogeneous Turbulence

Langevin Equation: Inhomogeneous Turbulence

How to determine unknown coefficients?

Fokker-Planck Equation

$$\frac{\partial P_E}{\partial t} + \frac{\partial u_i P_E}{\partial x_i} = -\frac{\partial (aP_E)}{\partial u_i} + \frac{1}{2} \frac{\partial^2 (b^2 P_E)}{\partial u_i^2}$$

Advection-diffusion for Eulerian velocity PDF – Eulerian equivalent of Langevin equation.

For derivation see: van Kampen, N.G.; 2nd ed., 1981. *Stochastic Processes in Physics and Chemistry*. North-Holland Pub. Co., 465 pp. Rodean, H. C., 1996: *Stochastic Lagrangian Models of Turbulent Diffusion*. Amer. Meteor. Soc., Boston, MA, 84 pp.

Solution in one dimension (unique):

$$\mathrm{d}w = \underbrace{\frac{1}{2} \frac{\partial \sigma_w^2}{\partial z} \mathrm{d}t}_{\mathbf{I}} - \underbrace{\left[\underbrace{\frac{C_0 \varepsilon}{2\sigma_w^2} - \frac{w}{2\sigma_w^2} \frac{\partial \sigma_w^2}{\partial z}}_{\mathbf{I}} \right] w \mathrm{d}t}_{\mathbf{I}} + \underbrace{(C_0 \varepsilon)^{1/2} \, \mathrm{d}\xi_i}_{\mathbf{III}}$$

I Drift correction term

- II Memory term
- III Diffusion term

Solution in three dimensions: method for determining Langevin coefficients is non-unique!

Thomson's (1987) 'simplest solution' (weak solution):

$$\mathrm{d} u_i = \frac{1}{2} \frac{\partial R_{il}}{\partial x_l} \mathrm{d} t - \frac{C_0 \varepsilon}{2} R_{ik}^{-1} u_k + \frac{1}{2} \frac{\mathrm{d} R_{il}}{\mathrm{d} t} R_{lj}^{-1} u_j \mathrm{d} t + (C_0 \varepsilon)^{1/2} \mathrm{d} \xi_i$$

 R_{ij} is the Reynolds stress tensor and R_{ij}^{-1} is its inverse

We can add any arbitrary rotation vector to the drift term and we'll still satisfy the well-mixed condition.

Langevin Equation: Rogue Trajectory Problem

$$\mathrm{d}w = \frac{1}{2} \frac{\partial \sigma_w^2}{\partial z} \mathrm{d}t - \left[\frac{C_0 \varepsilon}{2\sigma_w^2} - \frac{w}{2\sigma_w^2} \frac{\partial \sigma_w^2}{\partial z} \right] w \mathrm{d}t + (C_0 \varepsilon)^{1/2} \mathrm{d}\xi_i$$

It is possible for our Langevin equation to become unstable and get cases where $u_i \to \infty$

Langevin Equation: Rogue Trajectory Problem

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It is possible for our Langevin equation to become unstable and get cases where $u_i \to \infty$

ROGUE TRAJECTORY!

SSSSH! This is our dirty little secret.

What can we do about rogue trajectories?

- ad hoc constraints (violates well-mixed condition)
- Yee and Wilson (2007): semi-analytical scheme
- Postma et al. (2012): refine timestep
- Bailey et al. (2014): semi-implicit scheme

Langevin Equation: LES

Application to LES

$$\mathrm{d}w_s = \frac{1}{2} \frac{\partial \sigma_s^2}{\partial z} \mathrm{d}t - \left[\frac{C_0 \varepsilon_s}{2\sigma_s^2} - \frac{w_s}{2\sigma_s^2} \frac{\partial \sigma_s^2}{\partial z} \right] w_s \mathrm{d}t + (C_0 \varepsilon_s)^{1/2} \,\mathrm{d}\xi_i^*$$

Replace 'fluctuating' quantities with subgrid quantities

•
$$w \to w_s$$

•
$$\sigma^2 \rightarrow \sigma_s^2$$

• $\varepsilon \to \varepsilon_s$ (for Δ in inertial subrange, $\overline{\varepsilon} \approx \overline{\varepsilon_s} = -\tilde{S}_{ij} \tau_{ij}$)

*NOTE: this form assumes horizontal homogeneity and that τ_{ij} is isotropic. See Weil et al. (2004) for fully general version.

Langevin Equation: LES

e.g.,

- Kemp, J. R. and Thomson, D. J. (1996). Dispersion in stable boundary layers using large-eddy simulation. Atmos. Env. 30:2911-2923.
- Weil, J. C. and Sullivan, P. P. and Patton, E. G. (2004). The use of large-eddy simulations in Lagrangian particle dispersion models. *J. Atmos. Sci.* 61:2877-2997.
- Vinkovic, I., Aguirre, C., and Simoëns, S. (2006). Large-eddy simulation and Lagrangian stochastic modeling of passive scalar dispersion in a turbulent boundary layer. J. Turb. 7:N30.

End Current Literature (this is state-of-the-art)

LES Lagrangian Energy Spectra

No SGS model $(u_{si} = 0)$

LES Lagrangian Energy Spectra

No SGS model $(u_{si} = 0)$ SGS model

Where might all this energy be coming from?

- Langevin equation is inappropriate?
 Langevin coefficients are incorrect?
- 8 Rogue trajectories?

Where do ROGUE TRAJECTORIES come from?

Homogeneous version (1D RANS):

$$\mathrm{d}w = -\frac{C_0\varepsilon}{2\sigma^2}_{1/\tau_L} w\mathrm{d}t + (C_0\varepsilon)^{1/2} \,\mathrm{d}\xi$$

Where do ROGUE TRAJECTORIES come from?

Homogeneous version (1D RANS):

$$\mathrm{d}w = -\frac{C_0\varepsilon}{2\sigma^2} w\mathrm{d}t + (C_0\varepsilon)^{1/2} \,\mathrm{d}\xi$$

Inhomogeneous version (1D RANS):

$$\mathrm{d}w = \underbrace{\frac{1}{2} \frac{\partial \sigma^2}{\partial z}}_{\text{drift}} - \underbrace{\left[\frac{C_0 \varepsilon}{2\sigma^2} - \frac{w}{2\sigma^2} \frac{\partial \sigma^2}{\partial z}\right]}_{1/\tau_L} w \mathrm{d}t + (C_0 \varepsilon)^{1/2} \mathrm{d}\xi$$

Where do ROGUE TRAJECTORIES come from? Memory term:

$$-\underbrace{\left[\frac{C_0\varepsilon}{2\sigma_s^2} - \frac{w}{2\sigma_w^2}\frac{\partial\sigma_w^2}{\partial z}\right]}_{\frac{1}{\tau} = \frac{1}{\tau_1} - \frac{1}{\tau_2}} wdt$$

- au_1 : Local decorrelation time scale (isotropic)
- au_2 : Heterogeneity decorrelation time scale

$$-\underbrace{\left[\frac{C_0\varepsilon}{2\sigma_w^2} - \frac{w}{2\sigma_w^2}\frac{\partial\sigma_w^2}{\partial z}\right]}_{\frac{1}{\tau}}w\mathrm{d}t$$

What if τ turns out to be NEGATIVE? Or

$$C_0 \varepsilon < w \frac{\partial \sigma_w^2}{\partial z}$$

Recall our autocorrelation function:

$$\frac{\langle w(t)w(0)\rangle}{\langle w^2(0)\rangle} = e^{-t/\tau}$$

$$-\underbrace{\left[\frac{C_0\varepsilon}{2\sigma_w^2}-\frac{w}{2\sigma_w^2}\frac{\partial\sigma_w^2}{\partial z}\right]}_{\frac{1}{\tau}}w\mathrm{d}t$$

What could cause τ to be NEGATIVE?

 Δt not in the inertial subrange i.e., $\tau_L \lesssim \Delta t$

Thus $\frac{2\sigma_w^2}{C_0\varepsilon}$ is not the proper decorrelation timescale!

In this case, it is the problem not the discretization scheme that is unstable!!!!

Generalizing to 3D (assume τ_{ij} is isotropic)

$$\mathrm{d}u_{s,i} = \frac{1}{2} \frac{\partial \sigma_s^2}{\partial x_i} \mathrm{d}t - \underbrace{\left(\frac{C_0 \varepsilon_s}{2\sigma_s^2} - \frac{1}{2\sigma_s^2} \frac{\mathrm{d}\sigma_s^2}{\mathrm{d}t}\right)}_{1/\tau} u_{s,i} \mathrm{d}t + (C_0 \varepsilon_s)^{1/2} \mathrm{d}\xi_i$$

Generalizing to 3D (assume τ_{ij} is isotropic)

$$\mathrm{d}u_{s,i} = \frac{1}{2} \frac{\partial \sigma_s^2}{\partial x_i} \mathrm{d}t - \underbrace{\left(\frac{C_0 \varepsilon_s}{2\sigma_s^2} - \frac{1}{2\sigma_s^2} \frac{\mathrm{d}\sigma_s^2}{\mathrm{d}t}\right)}_{1/\tau} u_{s,i} \mathrm{d}t + (C_0 \varepsilon_s)^{1/2} \,\mathrm{d}\xi_i$$

Unstable if
$$C_0 \varepsilon_s < \frac{\mathrm{d}\sigma_s^2}{\mathrm{d}t}$$
 (this means τ is negative)

Generalizing to 3D (anisotropic τ_{ij})

$$\mathrm{d}u_{s,i} = \frac{1}{2} \frac{\partial \tau_{il}}{\partial x_l} \mathrm{d}t - \frac{C_0 \varepsilon_s}{2} \lambda_{ik} u_{s,k} + \frac{1}{2} \frac{\mathrm{d}\tau_{il}}{\mathrm{d}t} \lambda_{lj} u_{s,j} \mathrm{d}t + (C_0 \varepsilon_s)^{1/2} \mathrm{d}\xi_i$$

Generalizing to 3D (anisotropic τ_{ij})

$$\mathrm{d} u_{s,i} = \frac{1}{2} \frac{\partial \tau_{il}}{\partial x_l} \mathrm{d} t - \frac{C_0 \varepsilon_s}{2} \lambda_{ik} u_{s,k} + \frac{1}{2} \frac{\mathrm{d} \tau_{il}}{\mathrm{d} t} \lambda_{lj} u_{s,j} \mathrm{d} t + (C_0 \varepsilon_s)^{1/2} \, \mathrm{d} \xi_i$$

 $\begin{array}{l} \text{Unstable if} \\ G_{ij} = \delta_{ij} + \frac{\Delta t}{2} \left(-C_0 \varepsilon_s \lambda_{ij} + \frac{\mathrm{d}\tau_{il}}{\mathrm{d}t} \lambda_{lj} \right) \\ \hline \left| \lambda_{\max} \right| > 1 \end{array} \\ \left(\lambda_{\max} \text{ is largest eigenvalue of } G_{ij} \right) \end{array}$

Possible Solution: Reduce Δt

Sometimes not computationally feasible.

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Possible Solution: ad-hoc intervention

Violates well-mixed condition.

Possible Solution: Use mean quantities to calculate memory term

$$\mathrm{d}u_{s,i} = \frac{1}{2} \frac{\partial \sigma_s^2}{\partial x_i} \mathrm{d}t - \left[\frac{C_0 \overline{\varepsilon_s}}{2\overline{\sigma_s^2}} - \frac{u_{s,j}}{2\overline{\sigma_s^2}} \overline{\frac{\partial \sigma_s^2}{\partial x_j}} \right] u_{s,i} \mathrm{d}t + (C_0 \varepsilon_s)^{1/2} \, \mathrm{d}\xi_i$$

LES Energy Spectra

Possible Solution: Directly calculate $\tau_{L,s}$

We use Lagrangian scale-dependent SGS momentum model, which gives τ_{Ls}

See:

Stoll, R., and Porté-Agel, F. (2006). Dynamic Subgrid-Scale Models for Momentum and Scalar Fluxes in Large-Eddy Simulations of Neutrally Stratified Atmospheric Boundary Layers Over Heterogeneous Terrain. *Water Resour. Res.* 42:W01409.

$$\mathrm{d} u_{s,i} = \frac{1}{2} \frac{\partial \sigma_s^2}{\partial x_i} \mathrm{d} t - \frac{u_{s,i}}{\tau_{Ls}} \mathrm{d} t + (C_0 \varepsilon_s)^{1/2} \, \mathrm{d} \xi_i$$

this form is unconditionally stable!