

# Environmental Fluid Dynamics: Lecture 19

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# Stability Review

- In the broadest sense, stability is a dividing line between laminar and turbulent flows.
- If a flow is stable, it can become or remain laminar.
- If a flow is unstable, it can become or remain turbulent.
- There are many stabilizing and destabilizing factors to consider.



- These factors are often terms in the TKE balance equation
- To simplify and understand the problem, a destabilizing term is often paired with a stabilizing term to form a dimensionless ratio.
- This ratio is used to determine which factor "wins" and whether the flow becomes turbulent or not.
- Examples include the Reynolds number, Richardson number, Rossby number, Rayleigh number, and Froude number.



# Stability Review

- Recall that we previously discussed static stability.
- “Static” means no motion, so it described stability that is independent of the wind.
- Static stability determined whether a flow was capable of buoyant convection.
- Hence, we framed our discussion of static stability around vertical profiles of potential temperature.
- Today we will discuss dynamic instability.



# Dynamic Instability

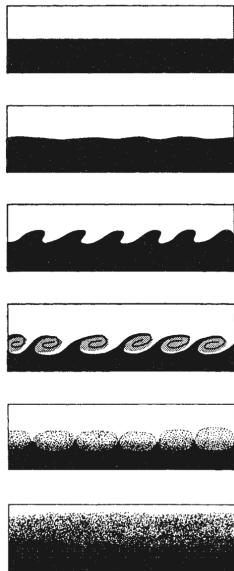
# Dynamic Instability

- “Dynamic” means having motion
- Unlike static stability, dynamic stability depends on the wind
- Even if the atmosphere is statically stable in some layer, wind shear may be sufficient to generate turbulence
- One example is Kelvin-Helmholtz instability





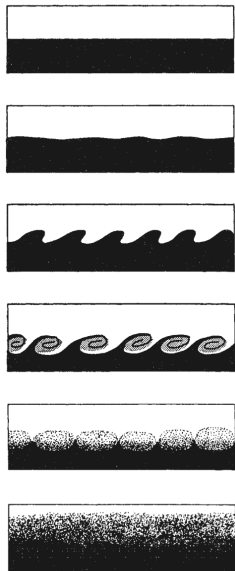
# Dynamic Instability: Kelvin-Helmholtz



- Consider low density fluid on top of high density fluid (statically stable)
- Imagine that shear ( $\Delta U$ ) exists at their interface.
- If shear becomes large enough, the flow is dynamically unstable,
- The amplitude of the waves grows until they break.
- The breaking wave is called a Kelvin-Helmholtz (KH) wave.
- The physics are different than a breaking wave on the ocean's surface, for instance.



# Dynamic Instability: Kelvin-Helmholtz



- Within each wave roll there are “packets” of statically unstable fluid (locally heavy above light fluid)
- The static and dynamic instabilities act to generate turbulence.
- The turbulence expands and leads to diffusion (mixing), which transfers momentum and reduces the shear.
- If the shear falls below the critical value, then the dynamic instability ends.
- In the absence of the shear, turbulence decays and the flow becomes laminar.



# Dynamic Instability: Kelvin-Helmholtz

- KH waves likely occur often in statically-stable shear layers, although they are rarely observed visually.
- If sufficient moisture is present in the atmosphere, clouds can form in the ascending portions of the wave.
- These are called billow clouds.



# Dynamic Instability: Kelvin-Helmholtz



photo by Paul Chartier, via: [Earthsky.org](http://Earthsky.org)

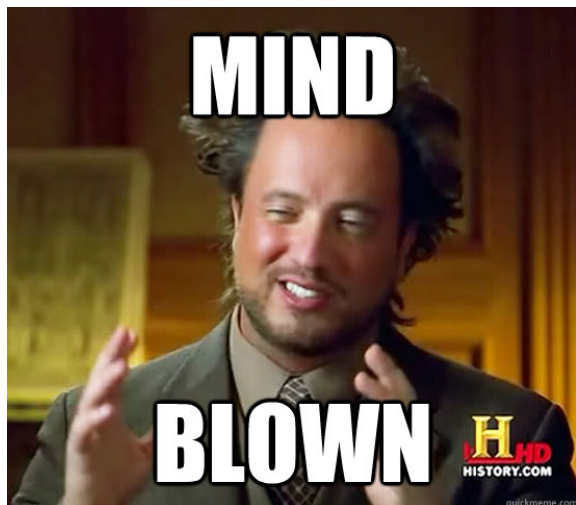


- Interestingly, both static and dynamic instabilities cause the flow to react in a way to remove the instability.
- So, turbulence is an "effort" by the flow to undo the cause of the instability.
- Static: convection moves buoyant air upward, which stabilizes the flow.
- Dynamic: turbulence reduces wind shears, which stabilizes the flow.



# Dynamic Instability

- In other words, turbulence exists to end its own existence



- Since observations show prolonged periods of turbulence in the atmosphere, there must be external forces that act to destabilize the PBL.
- Static: solar heating of ground.
- Dynamic: pressure gradients from synoptic-scale features causes winds to fight dissipation.



- To understand when a flow might become dynamically unstable, we can compare the relative magnitudes of the shear and buoyancy terms in our TKE balance equation.
- Shear leads to mechanical production of turbulence - a destabilizing effect.
- Buoyancy can either enhance or suppress turbulence - so it can be stabilizing or destabilizing.
- One such ratio is the Richardson number  $Ri$ .





# Richardson Number

# Dynamic Instability: Flux Richardson Number

- Consider a statically-stable environment.
- turbulent motions have to fight gravity, so buoyancy suppresses turbulence here.
- Conversely, wind shear acts to generate turbulence.
- The buoyant production/consumption term in the TKE balance is negative and the shear production term is positive.
- It is useful to examine their ratio.



# Dynamic Instability: Flux Richardson Number

- The ratio of buoyant production/consumption to shear production is called the **Flux Richardson Number**:

$$\text{Ri}_f = \frac{\left(\frac{g}{\theta_v}\right) (\overline{w'\theta'_v})}{(\overline{u'_i u'_j}) \frac{\partial \bar{u}_i}{\partial x_j}} = \frac{(\overline{w'b'})}{(\overline{u'_i u'_j}) \frac{\partial \bar{u}_i}{\partial x_j}}$$

- The Richardson number is dimensionless and contains nine terms in the denominator!



- If we neglect subsidence ( $\bar{w} = 0$ ) and assume horizontal homogeneity, we arrive:

$$\text{Ri}_f = \frac{\left(\frac{g}{\theta_v}\right) (\overline{w'\theta'_v})}{(\overline{u'w'}) \frac{\partial \bar{u}}{\partial z} + (\overline{v'w'}) \frac{\partial \bar{v}}{\partial z}} = \frac{(\overline{w'b'})}{(\overline{u'w'}) \frac{\partial \bar{u}}{\partial z} + (\overline{v'w'}) \frac{\partial \bar{v}}{\partial z}}$$



# Dynamic Instability: Flux Richardson Number

- For statically stable, unstable, neutral flows,  $Ri_f$  is positive, negative, 0 - respectively.
- Richardson suggested  $Ri_f = 1$  is a critical value since mechanical and buoyant terms are in balance.
- At any value less than  $Ri_f = 1$ , static stability is too weak to prevent mechanical generation of turbulence.
- For  $Ri_f < 1$ , buoyancy contributes to the generation of turbulence.

$$\left\{ \begin{array}{ll} \text{Flow is dynamically unstable (turbulent)} & Ri_f < 1 \\ \text{Flow is dynamically stable (laminar)} & Ri_f > 1 \end{array} \right.$$

- Note: statically unstable flow is by definition dynamically unstable.



# Dynamic Instability: Gradient Richardson Number

- With  $Ri_f$ , notice that our terms involve turbulent fluxes.
- So, it determines when a turbulent flow becomes laminar, but not when a laminar flow becomes turbulent.
- Using the idea of  $K$ -theory, we can say that  $-\overline{w'\theta'_v}$  is proportional to  $\partial\overline{\theta}_v/\partial z$ ,  $-\overline{u'w'}$  is proportional to  $\partial\overline{u}/\partial z$ , and  $-\overline{v'w'}$  is proportional to  $\partial\overline{v}/\partial z$ .
- We can use these assumptions to create a new Richardson number.



- The **Gradient Richardson Number** is given by:

$$\text{Ri} = \frac{\frac{g}{\theta_v} \left( \frac{\partial \bar{\theta}_v}{\partial z} \right)}{\left( \frac{\partial \bar{u}}{\partial z} \right)^2 + \left( \frac{\partial \bar{v}}{\partial z} \right)^2}$$

- If you see Ri without any subscript - the authors probably mean the gradient Richardson number.



- Theoretical and experimental data suggest that laminar flow becomes turbulent when  $Ri$  is below a critical value  $Ri_c$ , while turbulent flow ceases at some termination value  $Ri_t$ .

$$\begin{cases} \text{Laminar flow becomes turbulent} & Ri < Ri_c \\ \text{Turbulent flow becomes laminar} & Ri > Ri_t \end{cases}$$

- There is still debate about what values should be used, but typically  $Ri_c = 0.25$  and  $Ri_t = 1$ .
- There is a hysteresis since  $Ri_t > Ri_c$





## Why the hysteresis?

- The idea is that we need two conditions for turbulence: instability and a trigger mechanism.
- Consider KH waves as a trigger mechanism.
- In the absence of existing turbulence,  $Ri$  must fall well below  $Ri_t$  in order for KH waves to form.
- Experimental data suggests KH waves form when  $Ri < Ri_c$ .



## Why the hysteresis?

- Thus, we have a hysteresis.
- In other words, the Richardson number of a non-turbulent flow must be lowered to  $Ri_c$ , while a turbulent flow can remain so until the Richardson number surpasses  $Ri_t$ .



# Dynamic Instability: Bulk Richardson Number

- $Ri_c \approx 0.25$  is theoretically based on local gradients of wind and temperature.
- In the real world, we don't really know those local gradients.
- However, we can approximate them using discrete height layers.
- For example, we can approximate  $\partial\overline{\theta}_v/\partial z$  as  $\Delta\overline{\theta}_v/\Delta z$ .
- We can apply these approximations to the gradient Richardson number to create a new ratio.



- The **Bulk Richardson Number** is given by:

$$\text{Ri}_B = \frac{\frac{g \Delta \bar{\theta}_v}{\bar{\theta}_v \Delta z}}{\left(\frac{\Delta \bar{u}}{\Delta z}\right)^2 + \left(\frac{\Delta \bar{v}}{\Delta z}\right)^2} = \frac{g \Delta \bar{\theta}_v \Delta z}{\bar{\theta}_v [(\Delta \bar{u})^2 + (\Delta \bar{v})^2]}$$

- This is the form most often used by meteorologists.
- Observational and NWP data give wind and temperature at discrete points in the vertical.
- Note:  $\Delta \bar{u} = \bar{u}(\text{top}) - \bar{u}(\text{bottom})$



# Dynamic Instability: Bulk Richardson Number

- A few words of caution with  $Ri_B$  are warranted.
- The critical value is based on local gradients and not finite differences across layers.
- Gradients likely get “washed out” as the layer grows in thickness.
- Uncertainty arises in our ability to determine whether turbulence might form.
- Since some observations are spread far apart in the vertical, we must take care in interpreting the computed  $Ri_B$  when trying to characterize the considered flow.

