Environmental Fluid Dynamics: Lecture 18

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• We previously derived the following generic expression for turbulent momentum flux.

$$\begin{aligned} \frac{\partial(\overline{u'_ku'_i})}{\partial t} &= -\overline{u}_j \frac{\partial(\overline{u'_ku'_i})}{\partial x_j} - \left[\overline{u'_ju'_i} \frac{\partial\overline{u}_k}{\partial x_j} + \overline{u'_ku'_j} \frac{\partial\overline{u}_i}{\partial x_j} \right] - \frac{\partial(u'_ku'_ju'_i)}{\partial x_j} \\ &+ \overline{u'_kb'} \delta_{i3} + \overline{u'_ib'} \delta_{k3} \\ &- \left[\frac{\partial(\overline{u'_k\Pi'})}{\partial x_i} + \frac{\partial(\overline{u'_i\Pi'})}{\partial x_k} - \overline{\Pi'} \left(\frac{\partial u'_k}{\partial x_i} + \frac{\partial u'_i}{\partial x_k} \right) \right] \\ &+ \nu \frac{\partial^2(\overline{u'_ku'_i})}{\partial x_j^2} - 2\nu \frac{\overline{\partial u'_k}}{\partial x_j} \frac{\partial u'_i}{\partial x_j} \end{aligned}$$

 We will manipulate this equation to derive the turbulence kinetic energy (TKE) balance equation



A pedantic aside

- You might often see TKE written as *turbulent* kinetic energy.
- No! Stop that!
- Writing turbulent kinetic energy implies that there is some kinetic energy that is itself turbulent.
- The more accurate name is *turbulence* kinetic energy.
- Here, we describe kinetic energy that arises due to eddies in a turbulent flow.

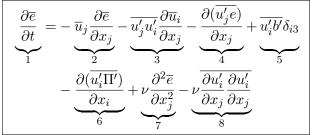


- We want an expression for TKE ($\overline{e} = 0.5\overline{u'_iu'_i}$).
- So, we set k = i and divide by 2:

$$\begin{split} \frac{1}{2} \frac{\partial (\overline{u'_i u'_i})}{\partial t} &= -\frac{1}{2} \overline{u}_j \frac{\partial (\overline{u'_i u'_i})}{\partial x_j} - \frac{1}{2} \left[\overline{u'_j u'_i} \frac{\partial \overline{u}_i}{\partial x_j} + \overline{u'_i u'_j} \frac{\partial \overline{u}_i}{\partial x_j} \right] - \frac{1}{2} \frac{\partial (u'_i u'_j u'_i)}{\partial x_j} \\ &+ \frac{1}{2} \overline{u'_i b'} \delta_{i3} + \frac{1}{2} \overline{u'_i b'} \delta_{k3} \\ &- \frac{1}{2} \left[\frac{\partial (\overline{u'_i \Pi'})}{\partial x_i} + \frac{\partial (\overline{u'_i \Pi'})}{\partial x_i} - \overline{\Pi'} \left(\frac{\partial u'_i}{\partial x_i} + \frac{\partial u'_i}{\partial x_i} \right) \right] \\ &+ \frac{1}{2} \nu \frac{\partial^2 (\overline{u'_i u'_i})}{\partial x_j^2} - \frac{1}{2} 2 \nu \overline{\frac{\partial u'_i}{\partial x_j} \frac{\partial u'_i}{\partial x_j}} \end{split}$$



• Substituting $\overline{e} = 0.5 \overline{u'_i u'_i}$ and simplifying yields



(1)



Terms in Eq. (1)

- Storage of tke
- 2 Advection of tke by the mean wind
- 8 Production of tke by the mean wind shear
- O Transport of the by turbulent motions (turbulent diffusion)
- 9 Production/destruction of tke by buoyancy
- **6** Transport of tke by pressure (pressure diffusion)
- Ø Molecular diffusion of tke
- 8 Viscous dissipation of tke



• $\partial \overline{e}/\partial t$ is very small at night, $\sim 2 \ m^2 \ s^{-2}$ during the day in the surface layer, and is often neglected over oceans.

- Advection can vary widely, but is usually considered negligible over large areas ($10 \text{ km} \times 10 \text{ km}$).
- The term becomes important over smaller areas, where heterogeneity matters.



- Shear production describes the interaction between the flux and gradient.
- The effect is strongest at the surface.

- Turbulent transport is advection by velocity fluctuations.
- This is not a source term because it integrates to zero over the entire domain (i.e., it redistributes TKE)



- Buoyancy generally maxes at during the daytime hours at $\sim 0.25~{\rm K~ms^{-1}}$
- Buoyancy is especially important during the day because it affects thermals.
- We can define the Deardorff scaling velocity for the mixed layer using vertical buoyancy flux.

$$w_* = (\overline{w'b'})^{1/3}$$

- For b > 0 (unstable), an air parcel displaced by turbulence continues to move in the direction of the displacement. Thus, there is production of
- For b < 0 (stable), an air parcel displaced by turbulence is forced to return to its starting point. Thus, there is destruction of TKE.



- Pressure fluctuations redistribute TKE.
- This is hard to measure and may be complicated by wave features.
- Accordingly, it is generally estimated as a residual.

- Molecular diffusion ranges (units of $m^2~s^{-3})$ from $\sim 10^{-11}$ in the ML to $\sim 10^{-7}$ in the SL
- The relatively small values mean that molecular effects are often neglected.



- The gradient $\partial u_i / \partial x_j$ is largest for small scales (i.e., shear is large for small-scale eddies).
- Viscous dissipation is always positive and ranges (units of $m^2~s^{-3})$ from $\sim 10^{-4}$ in the ML to $\sim 10^{-2}$ in the SL
- The more intense the small-scale turbulence, the stronger the dissipation.
- Small-scale turbulence is driven by large-scale turbulence via the energy cascade.
- TKE and dissipation usually follow each other closely.



Vertical profiles of various TKE budget terms

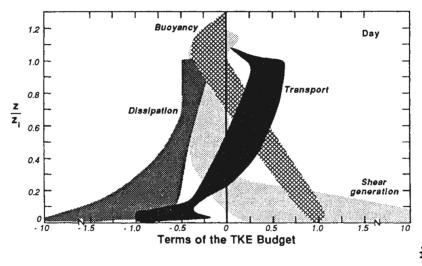


Fig. 5.4 from Stull (1988)

Buoyancy Variance Balance

• We previously derived the following generic expression for turbulent buoyancy flux.

$$\frac{\partial b'}{\partial t} + \frac{\partial}{\partial x_j} \left[\overline{u_j} b' + u'_j \overline{b} + u'_j b' - \overline{u'_j b'} \right] = -N^2 u'_j \delta_{j3} + \nu_h \frac{\partial^2 b'}{\partial x_j^2}$$

• We will manipulate this equation to derive the buoyancy variance balance equation



- We want an expression for buoyancy variance $(\overline{b'b'})$.
- So, we multiply by 2b' (2 so that we can invoke the product rule)

$$2b'\frac{\partial b'}{\partial t} + 2b'\frac{\partial}{\partial x_j}\left[\overline{u_j}b' + u'_j\overline{b} + u'_jb' - \overline{u'_jb'}\right] = -2b'N^2u'_j\delta_{j3} + 2b'\nu_h\frac{\partial^2 b'}{\partial x_j^2}$$



Buoyancy Variance Balance

• Make use of the product rule and then apply Reynolds averaging

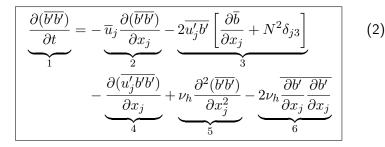
$$\frac{\partial(\overline{b'b'})}{\partial t} + \overline{u_j} \frac{\partial(\overline{b'b'})}{\partial x_j} + 2\overline{u'_j b'} \frac{\partial \overline{b}}{\partial x_j} + \frac{\partial(\overline{u'_j b'b'})}{\partial x_j} - 2\overline{b'} \frac{\partial \overline{u'_j b'}}{\partial x_j}$$
$$= -2N^2 \overline{u'_j b'} \delta_{j3} + \nu_h \frac{\partial^2(\overline{b'b'})}{\partial x_j^2} - 2\nu_h \frac{\partial \overline{b'}}{\partial x_j} \frac{\partial \overline{b'}}{\partial x_j}$$

• Simplification yields

$$\begin{aligned} \frac{\partial (\overline{b'b'})}{\partial t} + \overline{u_j} \frac{\partial (\overline{b'b'})}{\partial x_j} &= -2\overline{u'_jb'} \frac{\partial \overline{b}}{\partial x_j} - \frac{\partial (\overline{u'_jb'b'})}{\partial x_j} - -2N^2 \overline{u'_jb'} \delta_{j3} \\ &+ \nu_h \frac{\partial^2 (\overline{b'b'})}{\partial x_j^2} - 2\nu_h \frac{\partial \overline{b'}}{\partial x_j} \frac{\partial b'}{\partial x_j} \end{aligned}$$



• Grouping terms and simplifying yields



Terms in Eq. (2)

- Storage of buoyancy variance
- 2 Advection of buoyancy variance by the mean wind
- Ore Production of buoyancy variance by the mean buoyancy shear + stratification
- Transport of buoyancy variance by turbulence (turbulent diffusion)
- **5** Molecular diffusion of buoyancy variance
- **6** Viscous dissipation of buoyancy variance

